

## Hamon Potential Evapotranspiration

$$PET = k * 0.165 * 216.7 * N * \left( \frac{e_s}{T + 273.3} \right)$$

where,

PET	potential evapotranspiration [mm day <sup>-1</sup> ]
k	proportionality coefficient = 1 <sup>1</sup> [unitless]
N	daytime length [x/12 hours]
e <sub>s</sub>	saturation vapor pressure [mb]
T	average monthly temperature [°C]

### e<sub>s</sub> - saturation vapor pressure

$$e_s = 6.108e^{\left(\frac{17.27T}{T+237.3}\right)}$$

*Source: Lu et al. (2005)*

**script: calcPEThamon.r**

### Primary Sources

Allen et al. (1998). Crop evapotranspiration -- guidelines for computing crop water requirements. FAO Irrigation and Drainage Paper 56. United Nations, Rome.

Lu et al. (2005). A comparison of six potential evapotranspiration methods for regional use in the southeastern United States. Journal of the American Water Resources Association, 41, 621-633.

---

<sup>1</sup> Lu et al. (2005) uses a k value of 1.2 for the southeastern United States.

## **N - daylight hours in units of 12 hours**

$$N = \left(\frac{24}{\pi}\right) * \omega$$

where,

$\omega$  is the sunset hour angle [radians]

### **w - sunset hour angle**

$$\omega = \cos^{-1}[-\tan(\delta)\tan(\varphi)]$$

where,

$\varphi$  is latitude [radians]

$\delta$  is the declination [radians]

### **$\delta$ - declination**

$$d = 1 + 0.033\cos\left(\frac{2\pi}{365}J\right)$$

where,

J is the Julian Day of the year.

NB: when the sun does not rise  $\omega$  is set equal to 0, when the sun does not set  $\omega$  is set equal to  $\pi$ . This is accomplished by taking only the real portion of the result of the equation calculating  $\omega$ .

In order to calculate N at a monthly time step, we calculate average daily radiation for each day within the month and then average across the month.

*Source: Allen et al. (1998)*

**script: calcN.r**